

A complex network graph with numerous nodes (yellow and dark grey circles) connected by a dense web of black lines, representing a system of interconnected entities.

# Review october 2025

KARST GHOST-  
ROCK PROCESS

# Karsts

What is a karst ? :

- Environment
- Process
- System

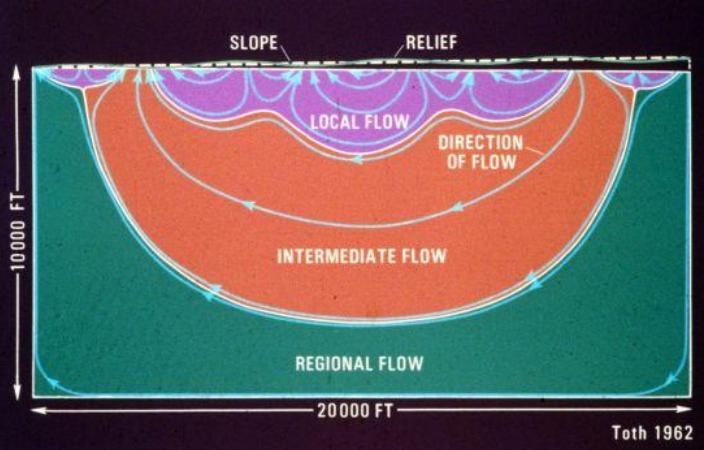
} Erosion of the host rock through physical and chemical processes by a fluid



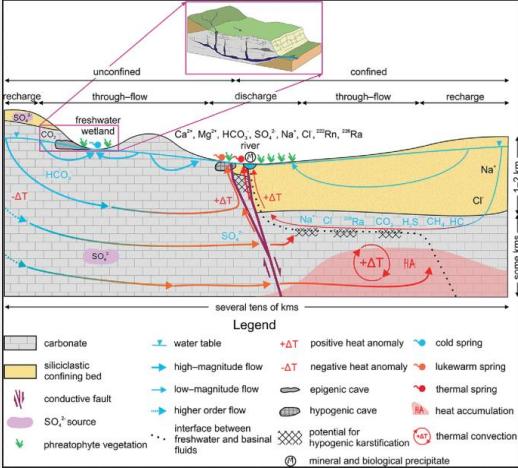
Broad variety of karsts :

- Endokarsts
- Exokarsts
- (Cryokarsts, paleokarsts, perched karst, etc...)

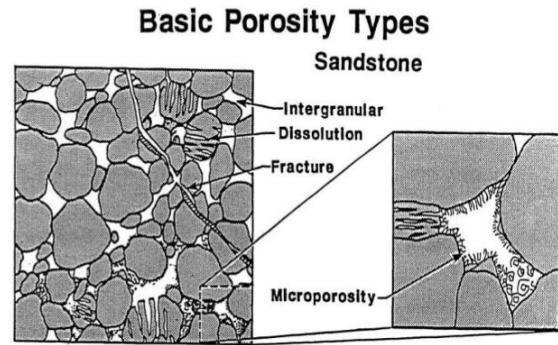
- Flow induced by a difference in the gradient pressure --> inlet (recharge area) and outlet (discharge area)
  - or : thermal gradient --> thermal buoyancy effect
- Fluid driven into the soils through different types of porosity --> connection of those porosities enhance the permeability of the rock
- Acidification mostly through dissolved atmospheric CO<sub>2</sub> enrichment of the fluid
- Dissolution of the minerals depends on the specific surface area --> area of the mineral which might be in contact with the fluid
- Flow following preferential paths : concentration of the advective alteration and dissolution until a breakthrough --> changing in the flow regime, concentration of the most part of the fluid flow in one main path
- 2 governing numbers in karst formation :
  - Péclet number : Describes the rate of influence between advection and diffusion processes in the model (depends on the fluid rate and the coefficient of diffusivity)
  - Damköhler number : Describes the rate of influence between reaction and flow of the reactives in the model
  - (a third indicative number, the Reynolds number, used in the calculation of the Péclet number, describes the linear or turbulent nature of the flow)



Tóth J., 1962

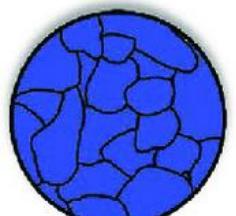


Mádlné Szőnyi J. et al., 2022



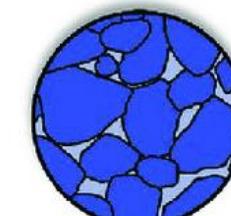
Radhi Wattan et al., 2022

### NO POROSITY



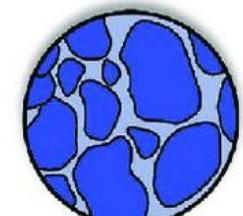
IMPERMEABLE

### CLOSED POROSITY



POROUS IMPERMEABLE

### INTERCONNECTED POROSITY



POROUS AND PERMEABLE

# Ghost-rock process



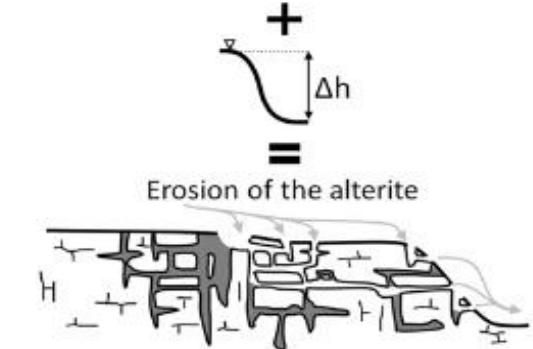
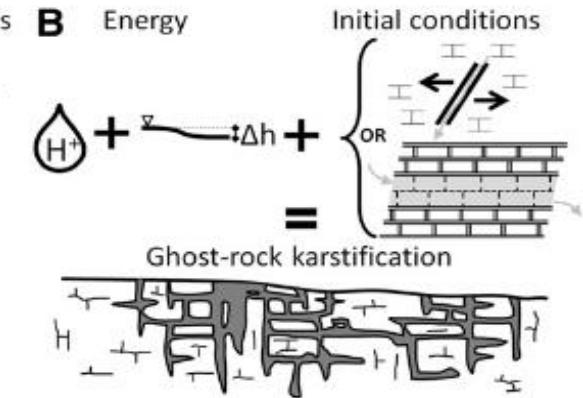
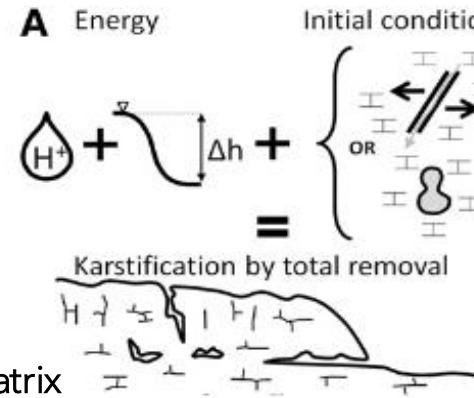
## I - Chemical alteration of the matrix

- Slow alteration of the matrix by a weak acidic fluid infiltrating in apertures such as fracturations
- High Damköhler number
- Low advection
- Dissolution affected by heterogeneities and tortuosity of the apertures



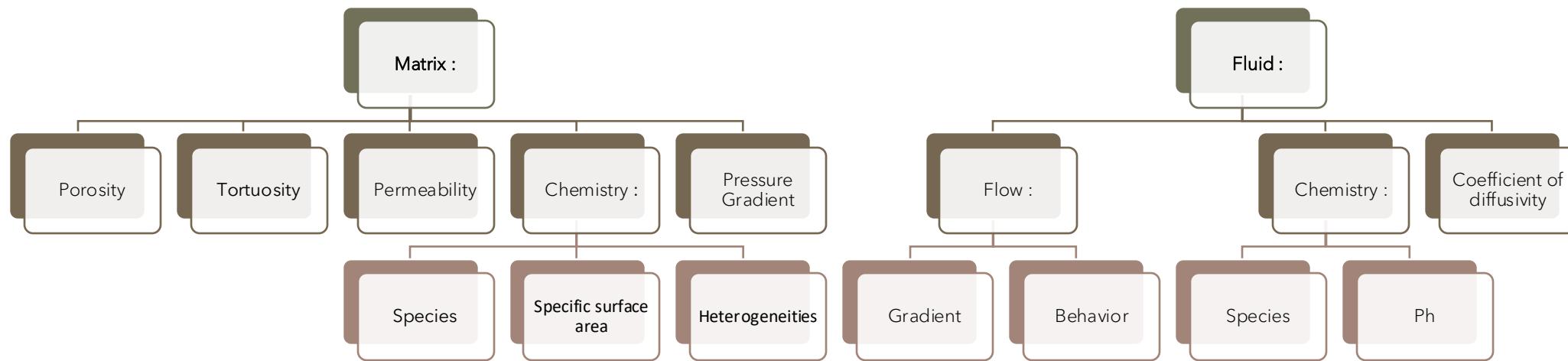
## II - Emptying the altered matrix

- Carried by advection of the fluid
- Collapsing by the effect of gravity
- Deep structures, emptied while getting closer to the surface, and affected by a changing in the hydraulic gradient



Dubois C. et al., 2014

# Key elements governing karst formation



# **Purpose of the simulation**

- Karsts : major water management possibilities : source of ¼ of water consumption worldwide, and major domestic water source in Mediterranean region (Ford D. et al.) --> need of more understanding of this phenomenon which can highly impact human water consumption
- Main objective of the simulation : numerical simulation of ghost-rock karstification process in a fractured matrix
- Simulations inspired by laboratory experiments carried-out by Linda LUQUOT in 2016 (Interaction between a fractured marl caprock and CO<sub>2</sub> -rich sulfate solution under supercritical CO<sub>2</sub> conditions) --> reproduction of this experiment on PFLOTRAN software

```
SIMULATION
  SIMULATION_TYPE SUBSURFACE
  PROCESS_MODELS
```

```
SUBSURFACE_FLOW flow
  MODE RICHARDS
  ! saturated flow
```

```
SUBSURFACE_TRANSPORT transport
  MODE GIRT
  OPTIONS
```

```
MULTIPLE_CONTINUUM
  /
```

```
/
```

```
SUBSURFACE
```

```
!=====
!===== numerical methods =====
```

```
  NUMERICAL_METHODS_FLOW
    LINEAR_SOLVER
    SOLVER DIRECT
  /
```

```
  NUMERICAL_METHODS_TRANSPORT
    LINEAR_SOLVER
    SOLVER DIRECT
  /
```

# Simulation

- Flow equation : Richards model (saturated matrix)
- Transport equation : GIRT (Global Implicit Reactive Transport) : requires high computational capacities --> OSRT (Operator-Split Reactive Transport) might be an alternative
- Solver : Linear (Newton ?)

$$\frac{\partial}{\partial t}(\phi s\eta) + \nabla \cdot (\eta \mathbf{q}) = Q_w, \quad (1)$$

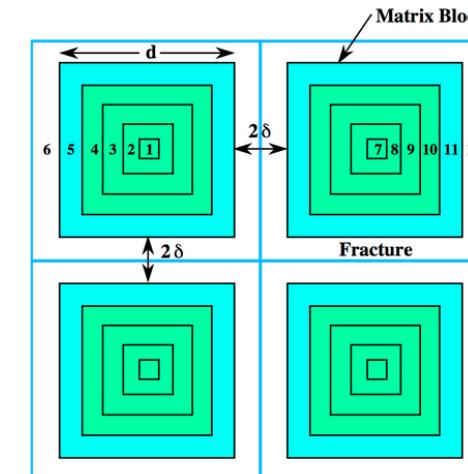
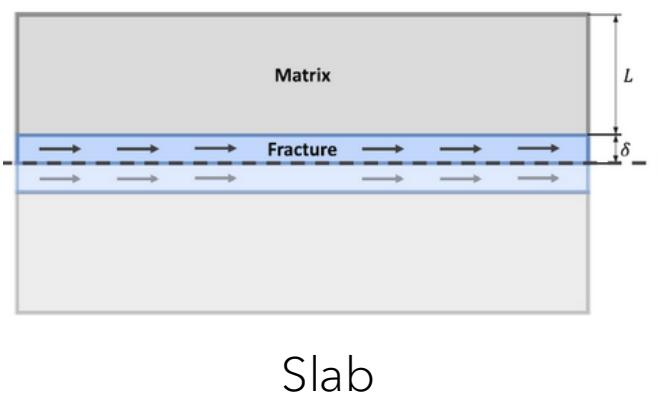
with Darcy flux  $\mathbf{q}$  defined as

$$\mathbf{q} = -\frac{kk_r(s)}{\mu} \nabla (P - \rho gz). \quad (2)$$

Here,  $\phi$  denotes porosity [-],  $s$  saturation [ $\text{m}^3 \text{ m}^{-3}$ ],  $\eta$  molar water density [ $\text{kmol m}^{-3}$ ],  $\rho$  mass water density [ $\text{kg m}^{-3}$ ],  $\mathbf{q}$  Darcy flux [ $\text{m s}^{-1}$ ],  $k$  intrinsic permeability [ $\text{m}^2$ ],  $k_r$  relative permeability [-],  $\mu$  viscosity [ $\text{Pa s}$ ],  $P$  pressure [ $\text{Pa}$ ],  $\mathbf{g}$  gravity [ $\text{m s}^{-2}$ ]. Supported relative permeability functions  $k_r$  for Richards' equation include van Genuchten, Books-Corey and Thomeer-Corey, while the saturation functions include Burdine and Mualem. Water density and viscosity are computed as a function of temperature and pressure through an equation of state for water. The source/sink term  $Q_w$  [ $\text{kmol m}^{-3} \text{ s}^{-1}$ ] has the form

# System Environment

- Fractured matrix
- High porosity (0.15) and permeability ( $1e-12$ )
- Secondary continuum : only diffusion considered



Nested cubes

```

=====
porous properties =====

CHARACTERISTIC_CURVES saturated ! For MODE RICHARDS. Relative permeability and saturation functi

      DEFAULT          ! Sets up dummy saturation and permeability functions for saturated sing
END

MATERIAL_PROPERTY porous
ID 1

      ! transient if CHEMISTRY/UPDATE_POROSITY
POROSITY 0.1500003d0          ! inhibé par CHEMISTRY/CALCULATE_INITIAL_POROSITY

      ! transient if CHEMISTRY/UPDATE_PERMEABILITY
PERMEABILITY
ISOTROPIC
PERM_ISO 1.d-12 m^2
/
PERMEABILITY_CRITICAL_POROSITY 0.0      ! theory guide : phi_c
PERMEABILITY_POWER 0.66667
PERMEABILITY_MIN_SCALE_FACTOR 0.0          ! theory guide : fmin

      ! transient if CHEMISTRY/UPDATE TORTUOSITY
TORTUOSITY 1.0
TORTUOSITY_POWER 1.0

LONGITUDINAL_DISPERSIVITY 0.0
TRANSVERSE_DISPERSIVITY_H 0.0
TRANSVERSE_DISPERSIVITY_V 0.0

CHARACTERISTIC_CURVES saturated ! For MODE RICHARDS. Relative permeability and saturation functi
SECONDARY_CONTINUUM
TYPE SLAB
LENGTH 0.01
NUM_CELLS 5
POROSITY 0.25d0
LIQUID_DIFFUSION_COEFFICIENT 5.1d-10
EPSILON 0.00909

END

```

```

MATERIAL_PROPERTY mat
ID 1

    ! transient if CHEMISTRY/UPDATE_POROSITY
    POROSITY 0.15000003d0      ! inhibé par CHEMISTRY/CALCULATE_INITIAL_POROSITY

    ! transient if CHEMISTRY/UPDATE_PERMEABILITY
    PERMEABILITY
        ISOTROPIC
        PERM_ISO 1d-12  m^2
    /                                         ! initial value for transient
    PERMEABILITY CRITICAL_POROSITY 0.0      ! theory guide : phi_c
    PERMEABILITY POWER 0.66667                ! theory guide : n=a doc 2015
    PERMEABILITY MIN SCALE FACTOR 0.0        ! theory guide : fmin

    ! transient if CHEMISTRY/UPDATE_TORTUOSITY
    TORTUOSITY 1.0
    TORTUOSITY_POWER 1.0                      ! theory guide : b

    LONGITUDINAL_DISPERSIVITY 0.0
    TRANSVERSE_DISPERSIVITY_H 0.0
    TRANSVERSE_DISPERSIVITY_V 0.0

    CHARACTERISTIC_CURVES saturated ! For MODE RICHARDS. Relative permeability and saturation functions and associated parameters.
END

MATERIAL_PROPERTY fracture
ID 2

    ! transient if CHEMISTRY/UPDATE_POROSITY
    POROSITY 1d0      ! inhibé par CHEMISTRY/CALCULATE_INITIAL_POROSITY

    ! transient if CHEMISTRY/UPDATE_PERMEABILITY
    PERMEABILITY
        ISOTROPIC
        PERM_ISO 8.333d-10  m^2
    /                                         ! initial value for transient
    PERMEABILITY CRITICAL_POROSITY 0.0      ! theory guide : phi_c
    PERMEABILITY POWER 0.66667                ! theory guide : n=a doc 2015
    PERMEABILITY MIN SCALE FACTOR 0.0        ! theory guide : fmin

    ! transient if CHEMISTRY/UPDATE_TORTUOSITY
    TORTUOSITY 1.0
    TORTUOSITY_POWER 1.0                      ! theory guide : b

    LONGITUDINAL_DISPERSIVITY 0.0
    TRANSVERSE_DISPERSIVITY_H 0.0
    TRANSVERSE_DISPERSIVITY_V 0.0

    CHARACTERISTIC_CURVES saturated ! For MODE RICHARDS. Relative permeability and saturation functions and associated parameters.
END

```

!===== fluid properties ======

FLUID\_PROPERTY

DIFFUSION\_COEFFICIENT 1.d-9 m<sup>2</sup>/s ! le

END

EOS WATER

DENSITY DEFAULT

! Internatio

VISCOSITY DEFAULT ! International Form

! DENSITY CONSTANT 997.16d0 kg/m<sup>3</sup>

! VISCOSITY CONSTANT 8.904156d-4 Pa·s

END

!\\ : Diffusion coefficient identical for all species in the simulation

!===== flow conditions ======

FLOW\_CONDITION initial\_pressure ! pour INITIAL\_CONDITION et BOUNDARY\_CONDITION outlet

! pour SUBSURFACE\_FLOW/MODE RICHARDS

TYPE

LIQUID\_PRESSURE HYDROSTATIC !.. DATUM, LIQUID\_PRESSURE, GRAVITY

/

DATUM 0.d0 0.d0 0.00015d0

LIQUID\_PRESSURE 201325.d0 Pa

Injected flux as inflow

END

FLOW\_CONDITION inlet flow

TYPE

LIQUID\_FLUX NEUMANN

! gradient constant

/

LIQUID\_FLUX 3.17d-5 m/s

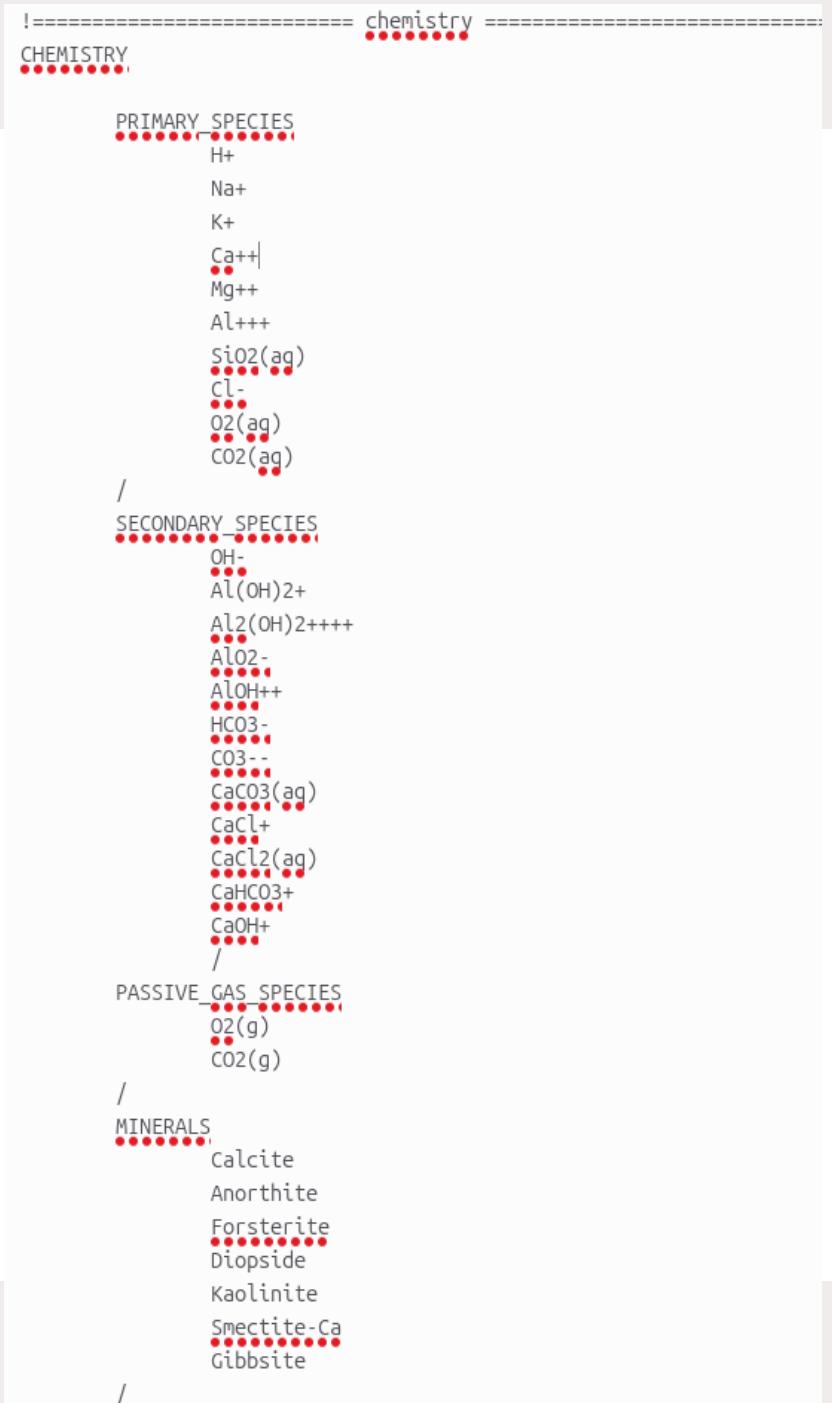
! vitesse Darcy constante

END

# Chemistry of the system

## Injection of a CO<sub>2</sub>-enriched water in a carbonate rock

- Weak acid --> high concentration but slow reaction
- 2 main processes of particles kinetics : diffusion and advection --> controlled by the Peclet number, requiring the diffusion coefficient (set as identical for every chemical species)
- Definition of primary species and secondary species to be precipitated
- Very low flow rate to be settled --> uniform dissolution ; if too high --> risk of wormhole structure, then breakthrough and end of the phantomisation



# Chemical mass conservation equations

ALPHA = 1 --> EQUATION  
SIMPLIFIED  
QJ --> SOURCE/SINK = 0

$$\frac{\partial}{\partial t} \left( \phi \sum_{\alpha} s_{\alpha} \Psi_j^{\alpha} \right) + \nabla \cdot \sum_{\alpha} \boldsymbol{\Omega}_j^{\alpha} = Q_j - \sum_m \nu_{jm} I_m - \frac{\partial S_j}{\partial t}, \quad (1)$$

and

$$\frac{\partial \phi_m}{\partial t} = \bar{V}_m I_m, \quad (2)$$

for minerals with molar volume  $\bar{V}_m$ , reaction rate  $I_m$  and volume fraction  $\phi_m$  referenced to an REV. The term involving  $S_j$  describes sorptive processes considered in more detail below. Sums over  $\alpha$  in Eqn. (1) are over all fluid phases in the system. The quantity  $\Psi_j^{\alpha}$  denotes the total concentration of the  $j$ th primary species  $\mathcal{A}_j^{\text{pri}}$  in the  $\alpha$ th fluid phase defined by

Lichtner P.C. et al., 2020

$$\Psi_j^{\alpha} = \delta_{\alpha} C_j^l + \sum_{i=1}^{N_{\text{sec}}} \nu_{ji}^{\alpha} C_i^{\alpha}, \quad (3)$$

In this equation the index  $l$  represents the aqueous electrolyte phase from which the primary species are chosen. The secondary species concentrations  $C_i^{\alpha}$  are obtained from mass action equations corresponding to equilibrium conditions of the reactions

Lichtner P.C. et al., 2020

$$\boldsymbol{\Omega}_j^{\alpha} = (\mathbf{q}_{\alpha} - \phi s_{\alpha} \mathbf{D}_{\alpha} \cdot \nabla) \Psi_j^{\alpha}. \quad (11)$$

The diffusion/dispersion tensor  $\mathbf{D}_{\alpha}$  may be different for different phases, e.g. an aqueous electrolyte solution or gas phase, but is assumed to be species independent. Dispersivity currently must be described through a diagonal dispersion tensor.

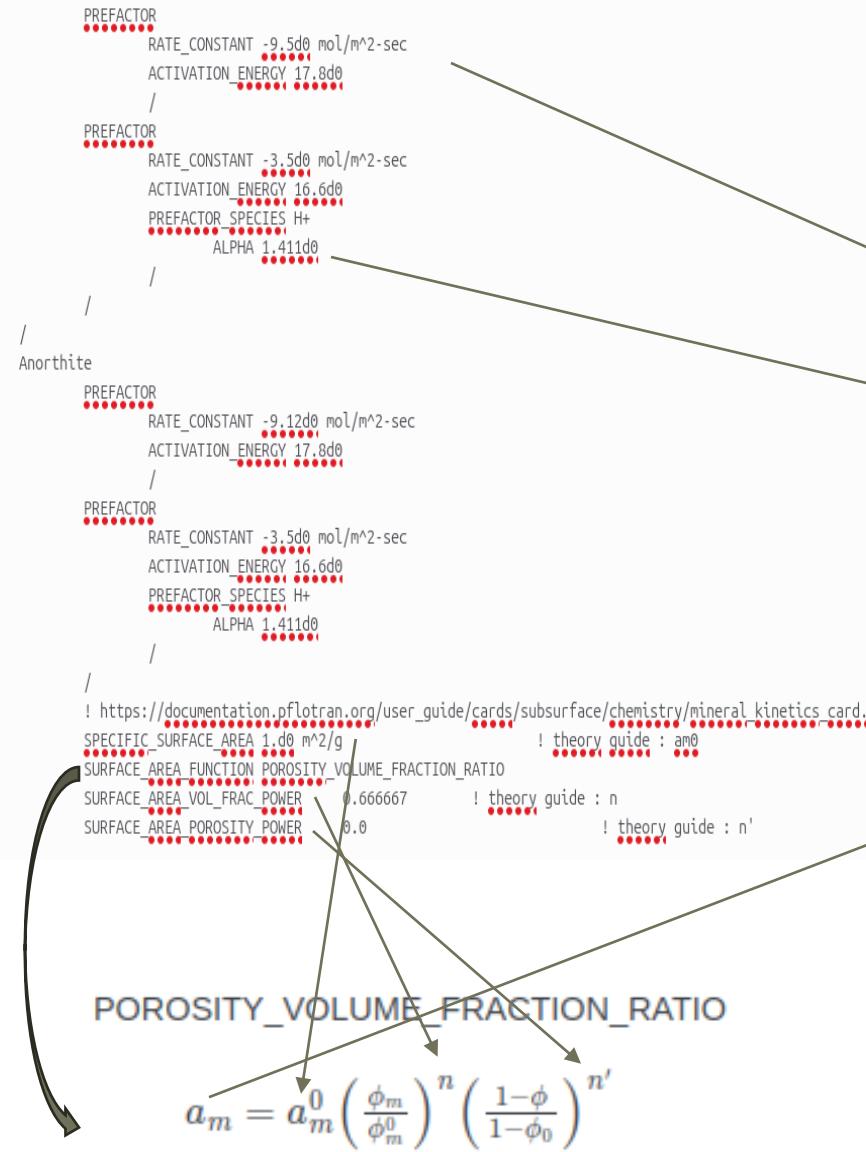
Lichtner P.C. et al., 2020

- linear  $K_D$  model:

$$S_j = \phi s_l K_j^D C_j = \hat{K}_j^D m_j,$$

with distribution coefficient  $\hat{K}_j^D$ .

Lichtner P.C. et al., 2020



Lichtner P.C. et al., 2020

$$\frac{\partial}{\partial t} (\phi \sum_{\alpha} s_{\alpha} \Psi_j^{\alpha}) + \nabla \cdot \sum_{\alpha} \Omega_j^{\alpha} = Q_j - \sum_m \nu_{jm} I_m - \frac{\partial S_j}{\partial t}, \quad (1)$$

and

$$\frac{\partial \phi_m}{\partial t} = \bar{V}_m I_m, \quad (2)$$

for minerals with molar volume  $\bar{V}_m$ , reaction rate  $I_m$  and volume fraction  $\phi_m$  referenced to an REV. The term involving  $S_j$  describes sorptive processes considered in more detail below. Sums over  $\alpha$  in Eqn. (1) are over all fluid phases in the system. The quantity  $\Psi_j^{\alpha}$  denotes the total concentration of the  $j$ th primary species  $\mathcal{A}_j^{\text{pri}}$  in the  $\alpha$ th fluid phase defined by

Lichtner P.C. et al., 2020

$$I_m = -a_m \left( \sum_l k_{ml}(T) \mathcal{P}_{ml} \right) \left| 1 - (K_m Q_m)^{\left( \frac{1}{\lambda_m \sigma_m} \right)} \right|^{\beta_m} \text{sign}(1 - K_m Q_m),$$

where a positive value corresponds to precipitation and a negative value to dissolution, and where

$a_m$  = mineral specific surface area [ $\text{m}^{-1}$ ]

$\mathcal{P}_{ml}$  = prefactor (a sum of prefactor rates; if activation energy is provided the Arrhenius equation is applied to each prefactor to calculate rates at different temperatures)

$K_m$  = equilibrium constant

$Q_m$  = ion activity product

$\sigma_m$  = Temkin number (default is 1)

$\lambda_m$  = mineral scaling factor (default is 1)

$\beta_m$  = affinity power (default is 1)

$k_{ml}$  = rate constant

Lichtner P.C. et al., 2020

===== chemical constraints =====

CONSTRAINT initial\_chem1 ! pour initial condition et boundary condition outlet

CONCENTRATIONS

! specie_name	concentration	constraint_type	constraint_species(optional)
H+	5.d0	Z	
Na+	5.0d-7	T	
K+	5.0d-7	T	
Ca++	4.0d-7	T	
Mg++	5.0d-7	T	
Al+++	5.0d-7	T	
SiO2(aq)	1.0d-7	T	
Cl-	5.0d-7	T	
O2(aq)	-0.699d0	G	O2(g)
CO2(aq)	-3.409d0	G	CO2(g)

/

MINERALS

! mineral_name	volume_fraction	specific_surface_area [m <sup>2</sup> mineral/m <sup>3</sup> bulk]
! Note : specific surface area supports units of area mineral per mass mineral [m <sup>2</sup> /g].		
Calcite	0.54	1.d0 m <sup>2</sup> /g
Anorthite	0.2	1.d0 m <sup>2</sup> /g
Forsterite	0.04	1.d0 m <sup>2</sup> /g
Diopside	0.07	1.d0 m <sup>2</sup> /g
Kaolinite	1.d-8	1.d0 m <sup>2</sup> /g
Smectite-Ca	1.d-8	1.d0 m <sup>2</sup> /g
Gibbsite	1.d-8	1.d0 m <sup>2</sup> /g

/

END

Constraint type options (Reactive Transport Mode):

- F = free ion concentration. Default units [mol/L]
- T = total aqueous component concentration. Default units [mol/L]
- P = pH
- PE = pe (for O<sub>2</sub>(aq) or H<sup>+</sup> only)
- M = concentration based on equilibrium with specified mineral. The float is an initial guess. Default units [mol/L]
- G = concentration based on equilibrium with a gas (partial pressure) [bars]
- L = Base 10 logarithm of concentration
- Z = charge balance
- TOTAL\_SORB = total sorbed concentration [mol/m<sup>3</sup><sub>bulk</sub>]
- TOTAL\_AQ\_PLUS\_SORB = total aqueous + total sorbed component concentration [mol/L]. The total sorbed concentration [mol/m<sup>3</sup><sub>bulk</sub>] must be converted to molarity [mol/L] by dividing by porosity \* saturation \* 1000 before being summed.

Lichtner P.C. et al., 2020

```

MINIMUM_POROSITY 0.0
CALCULATE_INITIAL_POROSITY
UPDATE_POROSITY

UPDATE_PERMEABILITY

UPDATE TORTUOSITY

DATABASE leo_database.dat

LOG_FORMULATION
ACTIVITY_COEFFICIENTS TIMESTEP ! Update after every time step.
! MOLAL

OUTPUT ! S'ajoute à OUTPUT/VARIABLES.
OFF
PH
TOTAL
ALL
MINERAL_SURFACE_AREA ! [m^2/m^3] ! Requires CHEMISTRY/OUTPUT/TOTAL
/
END

```



! inhibe MATERIAL\_PROPERTY/POROSITY

#### 0 - primary species and colloids fields:

- name
- a0 – ion size parameter
- Z – charge
- molarWeight

#### 1 - secondary species fields:

- name
- Number of species in aqueous complex
- For each species:
  - species stoichiometry
  - species name
- For each temperature:
  - logK
- a0
- Z
- molarWeight

#### 2 - gases fields:

- name
- molar volume (cm<sup>3</sup>/mol)
- Number of aqueous species in secondary reaction
- For each species:
  - species stoichiometry
  - species name
- For each temperature:
  - gas logK
  - gas molar weight

#### 3 - minerals fields:

- name
- molar volume
- Number of aqueous species in mineral reaction
- For each species:
  - species stoichiometry
  - species name
- For each temperature:
  - mineral logK
  - mineral molar weight

#### 4 - surface complexes:

- name
- Number of species in surface complexation reaction
- For each species:
  - species stoichiometry
  - species name

# Boundary conditions

- One inlet on the left border of the model : Flux of 3.17e-5 m/s
- One outlet at the right border of the model
- Other faces of the model : no specification --> defined as no flow

Type Dirichlet :  
Only advection considered -->  
no diffusion at the boundary conditions

```
! ===== transport conditions =====

TRANSPORT_CONDITION initial_chem1      ! pour INITIAL_CONDITION et BOUNDARY_CONDITION out
                                         TYPE ZERO_GRADIENT
                                         ! inflow : third-type or Robin boundary c

                                         CONSTRAINT_LIST
                                         0.d0 initial_chem1      ! appliqué à partir de t=0
                                         /
END

TRANSPORT_CONDITION outlet_chem         ! pour BOUNDARY_CONDITION inlet
                                         TYPE DIRICHLET_ZERO_GRADIENT
                                         ! inflow : constant concentration

                                         CONSTRAINT_LIST
                                         0.d0 inlet_chem1      ! appliqué à partir de t=0
                                         /
END

TRANSPORT_CONDITION inlet_chem         ! pour BOUNDARY_CONDITION inlet
                                         TYPE DIRICHLET_ZERO_GRADIENT
                                         ! inflow : constant concentration

                                         CONSTRAINT_LIST
                                         0.d0 inlet_chem ! appliqué à partir de t=0
                                         /
END

! ===== initial and boundary conditions =====

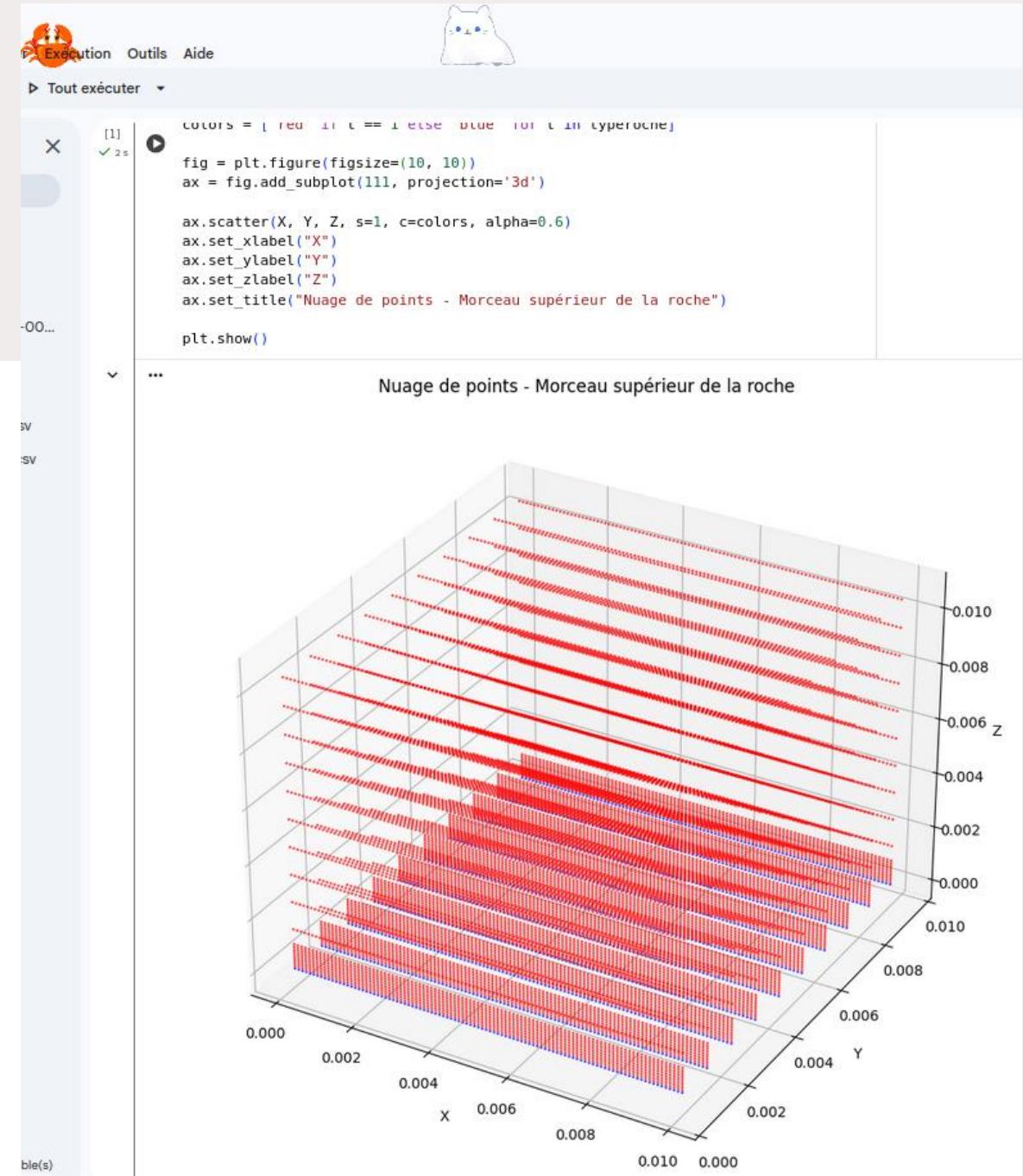
INITIAL_CONDITION
  FLOW_CONDITION      initial_pressure
  TRANSPORT_CONDITION initial_chem1
  REGION porous
END

BOUNDARY_CONDITION outlet
  FLOW_CONDITION      initial_pressure
  TRANSPORT_CONDITION outlet_chem
  REGION outlet
END

BOUNDARY_CONDITION inlet
  FLOW_CONDITION      initial_pressure
  TRANSPORT_CONDITION inlet_chem
  REGION inlet
END
```

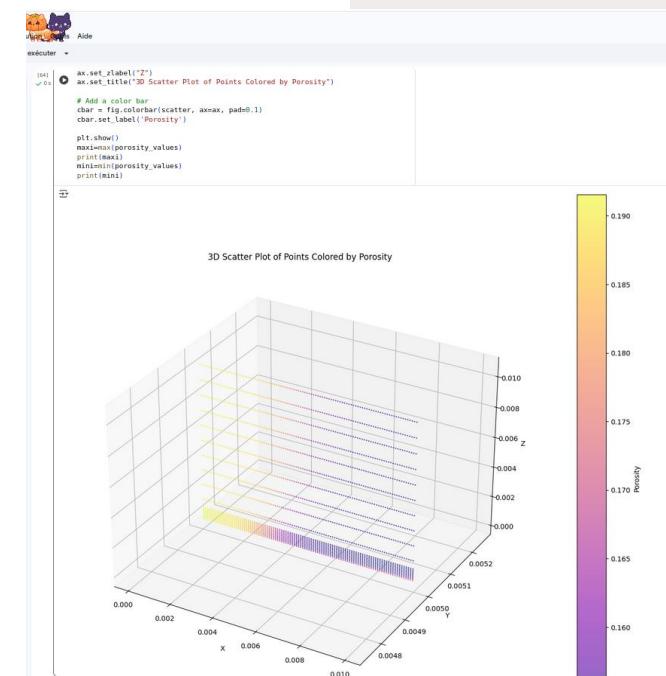
# Grid view of the model

- 1cm/1cm/1,1cm (X/Y/Z)
- On X axis :
  - 100 cells : 0.0001m resolution
- On Y axis :
  - 1 cell : 0.01m resolution (2D simulation)
- On Z axis :
  - 10 cells : 0.0001m resolution
  - 10 cells 0.001m resolution



# What was done so far

- Bibliography/Familiarisation with the topic
  - Karsts
  - Ghost rock
  - Geochemistry
- Introduction to PFLOTRAN
  - Test simulations with different grids/materials
  - Test simulations with a fracture in the model
- Modeling of a cloud points on python of the surface of a natural fracturation mesh

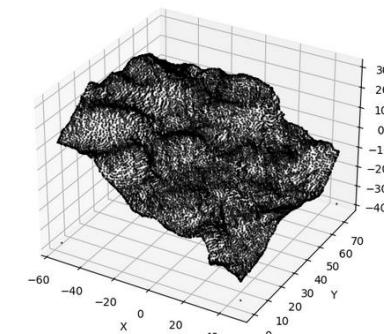


```
for ligne in f:
    x, y, z = map(float, ligne.strip().split())
    points.append([x, y, z])

X = [p[0] for p in points]
Y = [p[1] for p in points]
Z = [p[2] for p in points]

fig = plt.figure(figsize=(8, 6))
ax = fig.add_subplot(111, projection='3d')
ax.scatter(X, Z, Y, s=1, c='black', alpha=0.6)
ax.set_xlabel("X")
ax.set_ylabel("Y")
ax.set_zlabel("Z")
ax.set_title("Nuage de points - Morceau supérieur de la roche")
plt.show()
```

... Nuage de points - Morceau supérieur de la roche



# Next step

- Run a simulation reproducing the experiment of Linda :
  - New mineralogy
  - New fluid compositions and flow rates
  - New size of model
  - New matrix properties
  - New elements

**Table 1**  
Experimental conditions of the experiments.

Exp.	Injected solutions	Q [mLh <sup>-1</sup> ]	t [h]	h <sub>initial</sub> [m]	τ [s]	k <sub>f-initial</sub> [m <sup>2</sup> ]
1	<i>S-free</i> <i>I=0.3</i>	0.2	43	$3.2 \times 10^{-6}$	10.0	$8.8 \times 10^{-13}$
2	<b><i>S-rich I=0.6</i></b>	<b>0.2</b>	<b>35</b>	<b><math>3.8 \times 10^{-6}</math></b>	<b>12.0</b>	<b><math>1.2 \times 10^{-12}</math></b>
3	<i>S-free</i> <i>I=0.3</i>	1	21	$3.5 \times 10^{-6}$	2.0	$1.0 \times 10^{-12}$
4	<b><i>S-rich I=0.6</i></b>	<b>1</b>	<b>70</b>	<b><math>1.6 \times 10^{-5}</math></b>	<b>10.0</b>	<b><math>2.2 \times 10^{-11}</math></b>
5	<i>S-free</i> <i>I=0.6</i>	1	23	$4.1 \times 10^{-6}$ <sup>a</sup>	2.0	—
6	<i>S-free</i> <i>I=0.3</i>	60	6	$1.4 \times 10^{-5}$	0.1	$1.7 \times 10^{-11}$
7	<b><i>S-rich I=0.6</i></b>	<b>60</b>	<b>7</b>	<b><math>1.6 \times 10^{-5}</math></b>	<b>0.1</b>	<b><math>2.2 \times 10^{-11}</math></b>

t: experimental duration.

h<sub>initial</sub> and k<sub>f-initial</sub>: initial fracture aperture and fracture permeability, respectively.

τ: residence time.

<sup>a</sup> Permeability could not be measured. Aperture estimated from the ESEM image.

**Table 2**  
Chemical composition and saturation indexes of the initial synthetic solutions.

Solutions	S-free			S-free			S-rich			
	I [M]	0.3	0.6	0.6	4 × 10 <sup>-4</sup>	4 × 10 <sup>-4</sup>	4 × 10 <sup>-4</sup>	T [ °C]	25	25
Concentration (mol kgw <sup>-1</sup> )										
Ca	2.1 × 10 <sup>-2</sup>			4.2 × 10 <sup>-2</sup>			5.1 × 10 <sup>-2</sup>			
Mg	1.6 × 10 <sup>-2</sup>			3.2 × 10 <sup>-2</sup>			3.2 × 10 <sup>-2</sup>			
K	5.7 × 10 <sup>-3</sup>			1.1 × 10 <sup>-2</sup>			1.1 × 10 <sup>-2</sup>			
Si	4.5 × 10 <sup>-7</sup>			4.5 × 10 <sup>-7</sup>			4.5 × 10 <sup>-7</sup>			
S	1.0 × 10 <sup>-9</sup>			1.0 × 10 <sup>-9</sup>			2.7 × 10 <sup>-2</sup>			
Na	2.0 × 10 <sup>-1</sup>			3.6 × 10 <sup>-1</sup>			3.9 × 10 <sup>-1</sup>			
Fe	7.2 × 10 <sup>-9</sup>			7.2 × 10 <sup>-9</sup>			7.2 × 10 <sup>-9</sup>			
Al	1.9 × 10 <sup>-6</sup>			1.9 × 10 <sup>-6</sup>			1.9 × 10 <sup>-6</sup>			
Cl	2.7 × 10 <sup>-1</sup>			4.8 × 10 <sup>-1</sup>			5.1 × 10 <sup>-1</sup>			
Br	5.7 × 10 <sup>-3</sup>			1.1 × 10 <sup>-2</sup>			1.1 × 10 <sup>-2</sup>			
SI (Phreeqc)										
T [ °C]	25	60	25	60	25	60	25	60	25	60
pH (charge balance)	7.64	3.25	7.57	3.20	7.55	3.26				
pCO <sub>2</sub> [bar]	4 × 10 <sup>-4</sup>	61	4 × 10 <sup>-4</sup>	61	4 × 10 <sup>-4</sup>	61	4 × 10 <sup>-4</sup>	61		
Mineral										
Calcite	0.00	-3.60	0.00	-3.49	0.00	-3.32				
Dolomite	0.05	-6.97	0.07	-6.73	-0.02	-6.51				
Anhydrite	-7.80	-7.61	-7.72	-7.54	-0.28	-0.08				
Gypsum	-7.50	-7.66	-7.43	-7.59	0.01	-0.14				
Quartz	-2.34	-2.80	-2.30	-2.77	-0.29	-0.76				
Albite	-10.94	-20.23	-10.64	-20.26	-4.55	-14.24				
Clinochlore	-10.71	-55.23	-9.93	-55.07	-4.29	-49.05				
Illite	-14.08	-29.95	-13.70	-30.33	-6.68	-23.58				
Pyrite	-122.15	-217.9	-116.57	-60.75	-101.58	-205.3				
Kaolinite	-8.32	-17.85	-8.17	-18.36	-4.12	-14.62				

K<sub>eq</sub> values are from Phreeqc V.3 database except for gypsum which is from Garcia-Rios et al. (2014).

# Upcoming steps

- Continue to accumulate knowledge about karstic and Ghost rock mechanisms, as well as geochemistry and PFLOTRAN software
- Simulate in 2D, then in 3D the model with a fracture
- From this model :
  - Add heterogeneities to the model
  - Simulate different natures of matrix
  - Combine different fractures